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A lead isotope database for copper mineralization along the Guadalquivir River Valley and surrounding areas.

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Abstract:	<p>The Guadalquivir River Valley constitutes the main geographic environment for the earliest development of political structures in the South of the Iberian Peninsula. The inception of Cu-metallurgy over the millennia 4th to 2nd BC in this region coincides with the rise of social complexity as well as the strategic control of the territory based on the command of the supply of mineral resources. One of the main tools in the study of metallurgical processes and goods movement is the analysis of metal provenance by means of Pb isotope composition. This study includes 98 new Pb isotope analyses performed on Cu mineralizations, many of them with archaeological evidence of exploitation by the first metallurgical societies. The results provided here represent a substantial complement to those presented in published databases and allow for a better discrimination of potential sources for raw material supply of minerals used in the metallurgical processes. The existence of uncertainties introduced by the overlapping of Pb-isotope signatures of Cu-ores obtained from different geological contexts and presence of radiogenic Pb in a number of samples is pointed out.</p>	
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2 **A lead isotope database for copper mineralization along the Guadalquivir**
3 **River Valley and surrounding areas.**

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27 **Abstract**

1 28 The Guadalquivir River Valley constitutes the main geographic environment for the earliest
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3 29 development of political structures in the South of the Iberian Peninsula. The inception of
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5 30 Cu-metallurgy over the millennia 4th to 2nd BC in this region coincides with the rise of
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7 31 social complexity as well as the strategic control of the territory based on the command of
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9 32 the supply of mineral resources. One of the main tools in the study of metallurgical
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11 33 processes and goods movement is the analysis of metal provenance by means of Pb isotope
12
13 34 composition. This study includes 98 new Pb isotope analyses performed on Cu
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15 35 mineralizations, many of them with archaeological evidence of exploitation by the first
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17 36 metallurgical societies. The results provided here represent a substantial complement to those
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19 37 presented in published databases and allow for a better discrimination of potential sources
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21 38 for raw material supply of minerals used in the metallurgical processes. The existence of
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23 39 uncertainties introduced by the overlapping of Pb-isotope signatures of Cu-ores obtained
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25 40 from different geological contexts and presence of radiogenic Pb in a number of samples is
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27 41 pointed out.
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37 43 **Keywords:** Lead isotopes, ore deposits, Cu-metallurgy, database, Southern Iberian
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39 44 Peninsula, Spain, Portugal
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42 45
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57
58 52 of interest.
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53 **1. Introduction**

1 54

2
3 55 The Guadalquivir River valley (Spain) has been the main axis of communication throughout
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5
6 56 the history of the SW of the Iberian Peninsula and was crucial at forming the first political
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8 57 structure of regional extension in the prehistory of western Europe between the 4th and 2nd
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11 58 millennia B.C. (Nocete 2001, Nocete et al. 2010). Archaeological information indicates that
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13 59 the generalized production, movement and consumption of copper-based metallurgical
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16 60 products across this area were decisive in the development of a complex, hierarchical
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18 61 network of populated emplacements (Bayona 2008; Nocete 2001; Nocete et al., 2011). This
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21 62 social and geographical system included settlements that received metal products or that had
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23 63 a domestic type of productive activity, such as La Junta de los Ríos and Soto (Huelva) near
24
25 64 the Atlantic coast, and La Horca and Úbeda (Jaén) in the upper part of the valley (Nocete,
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27
28 65 2008; Nocete et al. 2010, 2011; Bayona, 2015, 2018). Also, large metallurgical centers have
29
30 66 been identified in the network. In some cases, they are linked to mining areas, as is the case
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33 67 of Cabezo Juré (Huelva) on the Iberian Pyrite Belt (Nocete 2004, 2006; Sáez et al., 2003).
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35 68 Other metallurgical centers, however, are located far from potential sources of raw material
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37
38 69 supply, such as Valencina de la Concepción and Carmona (Seville) at the former estuary of
39
40 70 the Guadalquivir, and Marroquíes (Jaén) in the upper valley (Nocete et al. 2008, 2011;
41
42
43 71 Bayona, 2015, 2018). The existence of centers of metallurgical production at a distance from
44
45 72 potential sources of minerals adds complexity to the process of creation of political
46
47
48 73 structures and calls for the need to evaluate their systems of harnessing resources. A first
49
50 74 step to this purpose requires identifying the potential sources of metal ores by means of a
51
52 75 reliable system.

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54 76 The PIGMALIOM Project was a regional archaeological study conducted to evaluate
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56
57 77 copper metallurgy and its role in the emergence and development of social complexity in the
58
59 78 Guadalquivir river valley (Nocete, 2000, 2001, Nocete et al., 2005 a, b; 2008, 2010, 2011).
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79 Systematic archaeological prospects and excavations in the Guadalquivir Basin during the
80 PIGMALION and related projects have revealed the recurring presence of copper artefacts
81 in all of the hundreds of settlements of the Third millennium BC investigated. Moreover, the
82 finding of copper ores, slags and crucibles in nearly thirty of the studied settlements indicate
83 a spread of copper metallurgy along the Guadalquivir Basin between the 4th and 2nd
84 millennia B.C. (Nocete, 2001, 2004, 2008; Nocete et al., 2008, 2010, 2011)

85 The sites of La Junta, Cabezo Juré and Soto in Huelva, Valencina de la Concepción
86 and Carmona in Seville, and Marroquíes, La Horca and Úbeda in Jaén (Fig. 1) are
87 representative of the different settlements along the axis of the most fertile soils and primary
88 copper mining areas. They cover the observed variations in population, economic
89 (settlements dedicated to mining, agriculture, etc.) and territorial variability (from 0.25 to
90 300 hectares of surface area). In these sites we developed a systematic program of
91 archaeological extensive documentation and radiocarbon dating of the direct contexts of
92 copper production (Nocete et al., 2011). From this program it was concluded that
93 metallurgical activity was generalized across the whole of the Guadalquivir Basin at the
94 beginning of the Third millennium BC. A complete understanding of this Chalcolithic
95 society needs, however, additional information about the supply areas and catchment
96 systems of copper ores resources, and about the circulation of copper ores and artefacts.
97 Appropriate correlations to determine the source areas and the transportation networks
98 require of good databases on the mineralizations located around the Guadalquivir Basin (Le
99 Guen et al. 1991; Stos-Gale et al. 1995; Klein et al. 2009; Stos-Gale and Gale 2009). Several
100 studies are available that report isotopic Pb data on metal mineralizations in the SW of the
101 Iberian Peninsula, basically conducted to understand the origin of the mineralizations and
102 their geodynamic framework (Marcoux 1998; Pomiès et al. 1998; García de Madinabeitia
103 2002; García de Madinabeitia et al. 2002; 2003, Tornos and Chiaradia 2004; Mateus et al.
104 2006). Furthermore, two interesting databases have been published to assist in the surveys

105 for the origin of archaeological ores (Santos Zalduegui et al. 2004; Klein et al. 2009). These
106 publications cover parts of the geographic field considered in our study (see Figures 1 and
107 2). In the aim to complete this information, we present 98 new analyses of Pb isotope
108 composition on selected samples of Cu mineralizations from the areas surrounding the
109 Guadalquivir Valley and its prolongation into south Portugal (Figure 2, Tables 1 and 2). A
110 summary of the type of work carried out in the field and of the investigated materials can be
111 seen in Figure 3.

112

113 2. Sampling and methods

114

115 Ore samples were obtained from mineralizations exposed on surface and from waste piles
116 after systematic prospection using the 1:50.000 geological maps of the Spanish Geological
117 Survey MAGNA series and their equivalents for south Portugal. Geographical location of
118 samples and mine names are indicated in Tables 1 and 2. Also, a kmz file from Google Earth
119 is provided as supplementary information for an easy tracking of the studied sample
120 locations.

121 The analytical results presented below have been obtained in the SGIker-
122 Geochronology and Isotope Geochemistry Facility of the University of the Basque Country
123 (Spain). Samples of copper ores were crushed in an agate mortar and dissolved with
124 concentrated nitric acid, and Pb was isolated from the Cu-rich matrix by conventional ion-
125 exchange chromatography in HBr-HCl media. All employed reagents were triple-distilled at
126 sub-boiling temperature before use. Lead isotopic ratios of samples labelled 'B', 'MRB',
127 'EAT' and "MIDAS 154" to "MIDAS 250" have been measured with a Finnigan MAT 262
128 thermal ionization mass spectrometer (TIMS). The measured ratios, given as an average of
129 10 blocks of 10 scans, were obtained in static collection mode using Faraday cups. The data
130 were corrected off-line for mass fractionation by comparison with replicate analyses of the

131 NBS-981 Pb standard. Individual errors in Cu ore analyses are similar to those obtained for
132 individual analysis of the NBS-981 standard (further details on the methodology are
133 provided in Santos Zalduegui *et al.* 2004). Lead isotope ratios for samples labelled 'MIDAS
134 359' to "MIDAS 489" have been obtained at the same facility using a Neptune multi
135 collector inductively coupled plasma mass spectrometer (MC-ICP-MS) with procedures
136 described in Rodríguez *et al.* (2020). All Pb isotopic data are presented in Table 2.

137

138 **3. Geological context of studied mineralizations**

139

140 During its upper and medium course, the Guadalquivir River flows along the limit between a
141 Variscan – pre-Variscan realm, and the Alpine Betic ranges of the southern Iberian
142 Peninsula. This stretch is occupied by the geological units that conform the Guadalquivir
143 foreland basin, bearing a triangular shape open to the Gulf of Cádiz (Figs. 1 and 2). This is
144 an ENE-WSW-trending 400 km long foreland basin developed during the late Miocene in
145 relation to tied flexural subsidence and stacking of thrust units of the external Betic
146 Cordillera (Barnolas *et al.*, 2019). The sedimentary infill is exposed along a topographic
147 gradient from the Sierra de Cazorla uplands to the Atlantic Ocean, where sedimentation
148 continues offshore in the abyssal plains of the Gulf of Cádiz. Along its northern boundary, a
149 rectilinear unconformable contact may be delineated over mostly Palaeozoic basement rocks
150 of the Iberian Massif. The Alpine orogenic domain to the South of the basin comprises
151 detritic and carbonate materials ranging in age from the Triassic to the Quaternary. Small
152 occurrences of mafic volcanic and subvolcanic rocks appear collated in the sedimentary
153 sequences. In the basement along the domain to the North of the foreland basin, a Variscan
154 and pre-Variscan orogenic region comprises rocks of Upper Proterozoic to Permian age that
155 belong to the South Portuguese (SPZ), Ossa-Morena (OMZ) and Central Iberian (CIZ) zones
156 of the Iberian Massif (Julivert *et al.* 1972; Quesada and Oliveira, 2019, 2020) (Fig. 2). The

157 lithologies are highly diverse although with a marked predominance of magmatic and
158 metamorphic types over those of a purely sedimentary origin. The regional structure
159 corresponds to a series of bands running in a WNW-ESE direction and essentially parallel to
160 the dominant directions arising from the Variscan orogeny. In many cases, these bands
161 represent units with differentiated geological characteristics. From the geographical
162 standpoint, it comprises a fairly homogenous morphological unit known as Sierra Morena.
163 Its overall relief is of the Appalachian valley and ridge type, with a marked Alpine
164 rejuvenation in the proximities of the river.

165 The main sources of supply of raw mineral materials used in the extractive
166 metallurgy of Cu would be related to various types of mineralizations associated with the
167 Variscan – pre-Variscan domain and, in the case of the Alpine domain, with Permo-Triassic
168 formations of red sandstones (New Red Sandstone), the development of which is limited to
169 the upper course of the Guadalquivir in Spain and to the so-called Mesozoic Orla in the
170 Algarve, Portugal.

171

172 3.1. *Mineralizations associated with the Variscan – pre-Variscan domain*

173

174 As noted earlier, this domain in the proximities of the Guadalquivir Valley has considerable
175 geological complexity, a fact that is also expressed in a great diversity of mineralizations
176 susceptible to have been sources of supply for ancient metallurgical activities.

177 The trait of the Valley from the NE to the SW successively traverses the southern
178 concealed prolongation of the three geological zones referred to above: CIZ, OMZ and SPZ
179 (see Fig. 2). The limits between zones have been established in terms of tectonic and
180 stratigraphic criteria (Julivert *et al.* 1972; Vera, 2004; Quesada *et al.*, 2019, 2020). Although
181 the precise location of these limits does not appear as being definitively closed, for the
182 purposes of this work, it is assumed that the limit between CIZ and OMZ is located in the

183 area known as Badajoz-Cordoba shear band (Burg *et al.* 1981; Ábalos *et al.* 1991a; Azor *et*
184 *al.* 1994) for the Variscan events and north of it, along the so-called Obejo-Valsequillo-
185 Puebla de la Reina domain, i.e. close to the Los Pedroches batholith, for the pre-Variscan
186 evolution (e.g. Eguíluz *et al.* 2000; Bandrés *et al.* 2004).

187 The limit between the OMZ and SPZ zones is better established and is considered to
188 be located in the South Iberian Shear Zone (Crespo-Blanc and Orozco 1988; Ábalos *et al.*
189 1991b; Simancas 2004; Ribeiro *et al.* 2010).

191 3.1.1. *The Central-Iberian Zone (CIZ)*

192
193 This zone is represented in the area studied by the Los Pedroches batholith and its
194 surroundings, and by the south-eastern foothills of the Alcuía Valley. The overall direction
195 of structures is WNW-ESE, caving obliquely with respect to the course of the Guadalquivir
196 at an approximate angle of 30°. In the proximities of the Guadalquivir River there are rock
197 outcrops with ages ranging from the Upper Proterozoic to the Carboniferous. Different types
198 of variably metamorphosed metasediments and magmatic rocks predominate. Igneous rocks
199 comprise an important part of the geological record of the region and have an important
200 effect on the metallogenic activity. Roughly speaking, two magmatic cycles are
201 distinguished: (i) a pre-Variscan one represented by volcanic and sub-volcanic rocks
202 interbedded in the Palaeozoic series (e.g. Palero 1991); and (ii) a late-Variscan cycle
203 represented in the study area by the Los Pedroches batholith and associated plutons
204 (Defalque *et al.* 1992; Larrea *et al.* 2004). The latter accounts for a complex magmatic
205 alignment that includes rocks with ages ranging from ca. 314 to 295 Ma (Bea 2004;
206 Carracedo *et al.* 2009). Associated mineralizations within the batholith and in adjacent areas
207 appear in fractures with different orientations. Mineralizations slotted into lengthwise
208 fractures and transversal to the predominant direction of the batholith correspond to the vein

209 type. Mineralizations of Cu and of Cu-U prevail within the batholith, whereas towards the
1 210 edges and in the peri-batholithic surrounding, mineralizations of Sn-W, Pb-Zn and of Bi are
2
3 211 more common. Towards the SE, the batholith dismembers into several individualized
4
5 212 granitic plutons (Cardeña, Linares, Guadalén/Arquillos) and the Santa Elena stock which
6
7 213 contain outstanding mineralizations of Pb-Zn-Ag and, to a lesser extent, Cu (Linares-La
8
9 214 Carolina mining areas). Towards the N of the batholith and outside its area of influence, lie
10
11 215 the mining districts of Almadén (Hg) and Valle de Alcudia (Pb-Zn ±Cu), that are not
12
13 216 included in this article as there are relatively recent works on the isotopic composition of Pb
14
15 217 for most of the sites already published (García de Madinabeitia 2002; Santos Zalduegui *et*
16
17 218 *al.* 2004; Higuera *et al.* 2005) and because they are quite distant from the area covered by
18
19 219 this study.
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25 220 In this work two sampling areas have been considered, one directly related with
26
27 221 different types of igneous rocks of the Los Pedroches batholith and associated plutons, cited
28
29 222 in Table 2 as “Los Pedroches batholith”, and other related to mineralizations in the nearby
30
31 223 host rocks, North of the batholith except sample MIDAS 233 to the South, and cited in
32
33 224 Table 2 as “Culm Pedroches”.
34
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37 225

38 226 3.1.2. *The Ossa-Morena Zone (OMZ)*

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41

42 228 The Ossa-Morena Zone is part of the southern branch of the Variscan orogen in the Iberian
43
44 229 Peninsula and was sutured to the Central Iberian and South Portuguese zones during
45
46 230 Cadomian and Variscan orogenic cycles (Eguíluz *et al.* 2000) (Fig.2). The dominant
47
48 231 structures trend NW-SE and, as for the CIZ, collide obliquely into the structural alignment
49
50 232 that sets the limit with the Alpine domain (*cif.*: Guadalquivir Fault). On a large scale, it
51
52 233 represents a mobile strike-slip zone with considerable tectonic, stratigraphic and
53
54 234 metallogenetic complexity (Tornos *et al.* 2002; 2004). Outcropping materials include
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235 diverse types of Precambrian and Palaeozoic series, as well as a great variety of magmatic
1 236 and low- to high-grade metamorphic rocks.
2

3 237 Due to the high geological complexity of the OMZ, different subzones, domains or
4
5 238 terranes have been proposed (e.g. more than 20 in Dallmeyer and Martínez García, 1990). In
6
7 239 this work, we have adopted a simplified scheme following more recent studies (e.g., Eguiluz
8
9 240 et al., 2000; Ábalos et al., 2002), which is as follows: (i) northern Ossa-Morena, related to a
10
11 241 Cadomian suture zone reactivated during Variscan times, it is bounded by the Peraleda fault
12
13 242 to the north and the Malcocinado fault to the south and includes the Obejo-Valsequillo
14
15 243 sector and the Blastomylonitic shear band; (ii) central Ossa-Morena, this is the largest sector
16
17 244 comprising most of the ‘domains’ of the literature (Olivenza-Monesterio, Barrancos-
18
19 245 Hinojales, etc.), with abundant rocks of the Ediacaran Serie Negra (Black Series), and
20
21 246 metamorphic areas and intrusives of Cadomian and Variscan ages, it goes from the
22
23 247 Malcocinado fault to the north up to the a series of minor faults to the south that separate it
24
25 248 from, (iii) southern Ossa-Morena, an area related to a Variscan suture that comprises the
26
27 249 metabasites of the Aracena-Acebuches metamorphic belt and a number of mafic to felsic
28
29 250 Variscan intrusives. For additional information and equivalence of domain names of the
30
31 251 literature with the simplified scheme adopted here, the reader is referred to the works
32
33 252 mentioned above.
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42 253 Unlike the surrounding areas (CIZ and SPZ), magmatism in the OMZ is
43
44 254 characterized by the abundance of basic rocks and by a continuity over time (Galindo and
45
46 255 Casquet 2004; Casquet and Galindo 2004) with peaks of high activity during the Upper
47
48 256 Proterozoic, Cambrian, Ordovician and Carboniferous. From the metallogenetic viewpoint,
49
50 257 the OMZ is characterized by the abundance of deposits associated with the Cadomian and
51
52 258 Variscan cycles (Tornos et al. 2004). The deposits associated with the Cadomian cycle show
53
54 259 general features typical of magmatic arcs along plate edges, whereas those from the
55
56 260 Variscan cycle cover a wide range of mineralization types, from magmatic to hydrothermal
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261 and sedimentary. Interpretation of a section of the middle crust imaged in seismic profiles as
1 262 a large body of mafic rocks related to the Variscan cycle (i.e., IRB: Iberseis Reflective
2
3 263 Body, Simancas et al. 2003) might support a link between the magmatic and metallogenetic
4
5 264 evolution of the OMZ during this orogenic cycle, though other interpretations also exist
6
7
8 265 (Pous et al., 2004; Puelles et al., 2014).

10 266 The distribution and diversity of mineralizations in the OMZ is remarkable. In their
11
12
13 267 synthesis on the metallogeny of the OMZ, Tornos et al. (2004) proposed the existence of a
14
15 268 series of metallogenetic belts which, roughly speaking, follow the same structural pattern as
16
17
18 269 the tectonostratigraphic domains proposed by Apalategui *et al.* (1990), although some of the
19
20 270 alignments proposed in fact overlap with the CIZ. Isotopic data for Pb from 26 different
21
22
23 271 types of mineralization by Tornos and Chiaradia (2004) show a broad range of compositions
24
25 272 and indicate two main metallogenetic events, associated with the Cadomian and Variscan
26
27
28 273 orogenic cycles. Most of the values presented by these authors fall below the growth curve
29
30 274 by Stacey and Kramers (1975), which suggests a mixture between crustal Pb and mantle
31
32
33 275 sources. This mixture appears to be more evident in the case of mineralizations associated
34
35 276 with the Variscan cycle (Tornos and Chiaradia, 2004).

37 277

40 278 3.1.3. *The South Portuguese Zone (SPZ)*

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45 280 The South Portuguese Zone occupies the SW part of the Iberian massif. The northern
46
47 281 boundary of the zone is marked by a band of metamorphosed mafic rocks, the Beja-
48
49
50 282 Acebuches amphibolites (Simancas 2004 and refs. included). Along the south side, it
51
52 283 borders with sediments of Alpine coverage in the Guadalquivir basin and its extension into
53
54
55 284 the Mesozoic sediments of the Portuguese Algarve (Barnolas et al., 2019). In the SPZ,
56
57 285 outcropping rocks show ages ranging from the Middle Devonian to the Permian.

286 Traditionally, three domains with different stratigraphic features have been
287 distinguished within the SPZ. The northernmost, defined as the Pulo do Lobo Group
288 (Schermerhorn 1971; Carvalho *et al.* 1971; Quesada, 1998) is made up of several
289 predominantly detritic formations that include mafic subvolcanic rocks with MORB affinity
290 (Giese and Bühn 1993) and whose ages have been established as Devonian (Oliveira 1990).

291 The central band of the SPZ is made up of the Iberian Pyrite Belt (IPB), which is
292 considered one of the metallogenetic provinces with the highest concentration of massive
293 sulphides in the world. The stratigraphic record of the IPB includes sedimentary and
294 magmatic rocks with ages ranging from the Upper Devonian to the Carboniferous. Its
295 evolution during Devonian times is marked by the sedimentation on a relatively shallow
296 siliciclastic platform (Moreno and Sáez 1990), giving rise to thick predominantly shale
297 sequences with sporadic intercalations of sandstones (quartz-arenites and quartz-rich wackes
298 of the so-called basal Phyllite-Quartzite Group or PQG). During the uppermost Devonian
299 and up to the Middle Viséan, an extensional episode occurred the consequences of which are
300 recorded in the fragmentation of the Devonian platform (Moreno *et al.* 1996), intense
301 bimodal magmatic activity (Sáez *et al.* 1996) and an important metallogenetic event of
302 massive sulphides that brought up the formation of the deposits of the IPB (Leistel *et al.*
303 1998; Sáez *et al.* 1999). This stage corresponds to the stratigraphic unit regionally known as
304 the Volcano Sedimentary Complex (VSC), which hosts the mineralizations of massive
305 sulphides characteristic of the region. Turbiditic synorogenic deposits of Culm facies overlie
306 the VSC in the region (Moreno 1993). During late- and post-Variscan times, a bimodal
307 intrusive event occurred in the SPZ. The most extensive outcroppings of these magmatic
308 rocks appear in the NE part of the Pyrite Belt comprising the batholith of the Sierra Norte of
309 Seville (De la Rosa and Castro 2004).

310 The southernmost domain of the SPZ is known as the Carrapateira Group or SW
311 Portuguese Domain. Its outcrops are circumscribed by the Aljezur and Bordeira antiforms.

312 The stratigraphic record covers a belt with ages similar to the IPB but, unlike this, there is no
1 313 magmatic activity associated with its evolution during the Devonian and the Carboniferous
2
3 314 periods. Data pointing to existence of metal mineralizations in this domain are lacking so
4
5
6 315 far.

8 316 As noted earlier, the mineralizations in the SPZ are concentrated in the Iberian Pyrite
9
10 317 Belt. There are two large groups of mineralizations in this domain: (a) pre-Variscan
11
12 318 mineralizations of Mn and massive sulphides related to volcanic-sedimentary processes, and
13
14 319 (b) vein-type mineralizations associated with the late- and post-Variscan evolution of the
15
16 320 region. The latter have received scarce coverage in the research studies probably due to its
17
18 321 small size and minor economic relevance. Nevertheless, like in the case of the massive
19
20 322 sulphides, many of them show evidence of exploitation in remote periods of history (Blanco
21
22 323 and Rothenberg 1981). The massive sulphides are syngenetic and occur associated with
23
24 324 volcanic and sedimentary rocks of the VSC. Isotopic data on Pb published to date for this
25
26 325 type of mineralization and related stockwork mineralizations (Marcoux 1998) highlight a
27
28 326 considerable degree of homogeneity at regional scale and also at the scale of the site
29
30 327 (Pomiès *et al.* 1998). Late- and post-Variscan hydrothermal mineralizations appear as vein
31
32 328 and small replacement masses after the sedimentary and volcanic rocks of the three main
33
34 329 lithological groups or units of the Belt (i.e., PQG, VSC and Culm), and also in veins of the
35
36 330 late-Variscan plutonic rocks. Generally, they show more radiogenic isotopic values of Pb
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38 331 than those common for the massive sulphides (Marcoux and Sáez 1994).

49 332 50 333 *3.2. Mineralizations associated with the Betic domain*

51 334
52 335 In inland zones of the Betic mountain ranges, there exist several ore deposits of widely
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54 336 diverse substances. However, as regards the influences on the area studied, the only
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56 337 recognized mineralizations of interest are found in the upper course of the Guadalquivir and
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338 in the Mesozoic Orla of the Algarve (Portugal). This entails “roller front” type
339 mineralizations of Cu associated with red sandstones dating from the Permian and Triassic
340 periods. They comprise small mineralized bodies with associations of minerals typical of the
341 oxidized zones of Cu sites, including malachite, azurite and cuprite as characteristic
342 minerals. The most important deposits are located in the surroundings of Navas de San Juan,
343 Jaén (Torres and Fernández 1983). Similar mineral associations have been identified as
344 related to red sandstones near Silves in the Algarve (Portugal).

346 4. Pb isotope results

347
348 The results of the isotopic analyses of Pb ores conducted for this study and the geographic
349 location by means of UTM coordinates of the sampled sites, together with their host-rock
350 environment, are shown in the Tables 1 and 2. Their location in the geological context of the
351 Iberian southwest is shown in Figure 2. The data obtained are presented in the diagrams
352 commonly used in the analysis of metal provenance as applied in archaeometallurgical
353 studies. These diagrams are $^{207}\text{Pb}/^{206}\text{Pb}$ versus $^{208}\text{Pb}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ versus $^{206}\text{Pb}/^{204}\text{Pb}$.
354 The general representation of the data (Fig. 4) highlights a wide range of values, out of
355 which three main groups, hereafter referred to as Groups 1, 2 and 3, have been established.
356 One of these groups, the Group 2, is in turn subdivided in a number of more detailed
357 subgroups. Their composition from less radiogenic to more radiogenic is as follows:

358 **Group 1:** It is represented by just two samples from the La Preciosa mine (Peñaflor,
359 Seville), whose mineralization, pyrite and chalcopyrite principally, is associated to Upper
360 Proterozoic rocks of the high-grade metamorphic core of Lora del Río (central OMZ).
361 Similar values have been reported for mineralizations elsewhere within old metamorphic or
362 sedimentary domains of the the Ossa-Morena Zone and also of the Alcuía Valley in the
363 Central Iberian Zone (Santos Zalduegui et al., 2004; Tornos and Chiaradia, 2004).

364 **Group 2:** This is the main group comprising 79 samples. It forms a compact cluster
1 365 covering mineralizations of almost all the geological contexts considered in this study. A
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3 366 close inspection expanding the scale of the axes highlights the existence of 4 main
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6 367 subgroups: 2A, 2B, 2C and 2D, while an additional enlargement of the diagram allows to
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8 368 subdivide further 2A and 2B into 2A1, 2A2, 2B1 and 2B2 subgroups (Figure 4).
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11 369 **Subgroup 2A1:** It includes 17 samples, out of which 16 are from the Iberian Pyrite
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13 370 Belt in the SW of the investigated area (SPZ) that correspond to: (i) 12 samples of massive
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15 371 sulphides, specially pyrite, chalcopyrite ± galena, and associated gossans (with goethite and
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18 372 hematite as representative minerals) related to the Carboniferous VSC; and, (ii) 4 samples
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20 373 with more complex parageneses from veins related to the Devonian materials of the PQG.
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22
23 374 The one exception within this group is sample MIDAS 437 from Posadas, within the
24
25 375 Cadomian Olivenza-Monesterio antiform in central OMZ.
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28 376 **Subgroup 2A2:** It forms the largest cluster within Group 2, with 35 points that
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30 377 correspond essentially to 28 samples from the Linares-La Carolina mining area (Jaén) and
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32 378 nearby localities (Guarromán, Carboneros, Baños de la Encina, Villanueva de la Reina,
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34
35 379 Andújar, Vilches, Cardeña and Montoro). In this case, the main mineral is galena with
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37 380 associated pyrite and chalcopyrite, while secondary phases are cerussite and malachite,
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40 381 among others. All of them are associated to Variscan granodiorites-monzogranites of the
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42 382 Los Pedroches batholith and the metamorphic rocks of the contact aureole or their
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45 383 sedimentary precursors, that is, the Carboniferous Culm of Los Pedroches in the Central
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47 384 Iberian Zone. This subgroup also includes a few scattered mines toward the West of this
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50 385 area, namely three samples related to metamorphic rocks of Cambrian age from the Cerro
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52 386 Muriano area in northern Ossa-Morena, one rare sample (MIDAS 248) close to the Cala ore
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54 387 district (La Vicaría mine) in central Ossa-Morena, which departs from the composition of
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57 388 other Cala samples analyzed in this work (see below). Finally, three samples are from veins
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389 in the Iberian Pyrite Belt, two of them (MRB158, MIDAS 448) from the Portuguese section
1 390 of the IPB.

2
3 391 **Subgroup 2B1:** It corresponds to samples from 8 mines; 4 from the Cardeña area
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5 392 within igneous rocks (granite porphyries) of the Los Pedroches batholith in the CIZ, 3 from
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7 393 scattered outcrops in the Cheles-Barrancos sector of central Ossa-Morena, also associated to
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9 394 igneous rocks, and one from a vein in the Iberian Pyrite Belt. In this case the parageneses are
10
11 395 more complex. In general, galena is very abundant with secondary minerals as malachite,
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13 396 cerussite, chrysocolla and goethite.

14
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16 397 **Subgroup 2B2:** It includes 11 mines and is the most heterogeneous subgroup in
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18 398 terms of sample provenance. Four samples are from veins in the Iberian Pyrite Belt. Other 4
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20 399 are from mine shafts at Hornachuelos in Central Ossa-Morena, geographically to the SW
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22 400 and close to the sampling area of Group 2B1. Two samples are from host rocks of the Los
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24 401 Pedroches batholith. And, finally, one sample is from the Obejo-Valsequillo domain in
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26 402 Northern Ossa-Morena. As in case of group 2B1, secondary minerals are very frequent,
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28 403 specially malachite.

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30 404 **Subgroup 2C:** It comprises samples from 5 mines that are distinctly more radiogenic
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32 405 than those of the previous sets. The subgroup is also heterogeneous in terms of areal
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34 406 distribution. Three samples are associated to the Permo-Triassic sedimentary materials of
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36 407 the New Red Sandstone in the Betic domain, one sample is from the Obejo-Valsequillo area
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38 408 of northern Ossa-Morena and another one is from a vein in the Iberian Pyrite Belt. As in the
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40 409 case of subgroups 2B1 and 2B2, secondary minerals (malachite, azurite) are notorious.

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42 410 **Subgroup 2D:** It is made of samples from 3 mines geographically scattered that, in
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44 411 the $^{207}\text{Pb}/^{206}\text{Pb}$ versus $^{208}\text{Pb}/^{206}\text{Pb}$ and $^{207}\text{Pb}/^{206}\text{Pb}$ versus $^{206}\text{Pb}/^{204}\text{Pb}$ diagrams, are related to
45
46 412 the samples of subgroup 2A but without a clear relationship with the previously defined
47
48 413 clusters. It includes sample MIDAS 441, from El Centenillo, the only sample in this work
49
50 414 collected in Ordovician materials from the Alcudia Anticline (CIZ).

415 **Group 3:** This group is characterized by highly radiogenic lead (Table 2 and Fig. 4),
1 416 and comprises ten samples from different localities (Andújar, Baños de la Encina and
2
3 417 Villanueva de la Reina) in the Los Pedroches batholith of the Central Iberian Zone. Three
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5 418 more samples are from the Cheles-Barrancos sector and two are from the Cala skarn, all in
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8 419 central Ossa-Morena. Finally, another three samples are associated with IPB veins from
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11 420 Portugal. At the Los Pedroches batholith, the samples correspond to Cu-U mineralizations,
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13 421 sometimes with presence of torbernite and autunite. In fact, the presence of exploitable U
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15 422 ores in this sector of the batholith has been known for long (e.g., Arribas Moreno, 1963,
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18 423 Ministerio de industria, 1971). Similar large dispersion of analytical data in Pb/Pb diagrams
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20 424 has been related to the presence of U-bearing minerals elsewhere within the Iberian massif
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23 425 (e.g., Huelga-Suárez, 2014).

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27 427 **5. Conclusions**

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33 429 The 98 new Pb isotope analyses on Cu mineralizations, many of them with
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35 430 archaeological evidence of exploitation by the first metallurgical societies, presented in
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37 431 Tables 1 and 2 and commented in previous sections of this study represent a substantial
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40 432 complement to those already available in published databases. These new data will allow
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42 433 for a better discrimination of potential sources for raw material supply of minerals used in
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44
45 434 the metallurgical processes. Our results, that concern principally the Pb isotope composition
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47 435 of Cu mineralizations from the areas surrounding the Guadalquivir Valley and its
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50 436 prolongation into south Portugal, clearly show that those samples from the groups where the
51
52 437 primary mineralization predominates, like those of subgroup 2A1 (sulfides or directly
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54 438 related gossans from the Iberian Pyrite Belt) or subgroup 2A2 (galenas from the Los
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56
57 439 Pedroches batholith area), present a restricted range of isotope values and plot on well-
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59 440 defined fields of the Pb/Pb diagrams. A note of caution is however imposed since there exist

441 some uncertainties due to the overlapping of Pb-isotope signatures of Cu-ores in provenance
1 442 of different geological contexts. Thus, the isotopic compositions of samples from groups in
2
3 443 which the secondary mineralization appears to have been extensive (subgroups 2B1 and
4
5 444 2B2), present a greater dispersion regardless of the geological domain or context to which
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7 445 they are associated. Finally, the deposits of the so-termed Group 3, in all probability
8
9 446 associated with minor U mineralizations, constitute a special case. Samples of this group
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11 447 show a large dispersion of isotope values albeit following a linear tendency, a fact that in
12
13 448 principle is to be linked to their U content and the radioactive disintegration of this element.
14
15 449 The identification of U-bearing copper ores that could have been exploited in ancient times
16
17 450 in the south of the Iberian peninsula may give a successful provenance interpretation to rare
18
19 451 copper and bronze artifacts from the region that have highly-radiogenic lead isotope ratios.
20
21 452 In our experience, it is not uncommon that one or some of these artifacts are analyzed
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23 453 together with a larger number of isotopically-conventional (i.e., less radiogenic)
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25 454 archaeological samples. Their oddity make them prone to be discarded in provenance
26
27 455 interpretations and their isotopic data are eventually not published.
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458

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466

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720 **Figure captions**

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3 722 **Figure 1.** Oblique perspective of the Guadalquivir Valley and its surroundings with its

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6 723 prehistoric estuary and the location of the 8 main settlements mentioned in the text

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11 725 **Figure 2.** Geological map of southern Iberia with the principal geological zones and the

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13 726 location of the analyzed ores. Areas covered by previous published lead isotope data

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15 727 bases are shown by dotted or dashed colored lines.

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20 729 **Figure 3.** Examples of the mines and materials studied in connection with the PIGMALION

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23 730 Project. A) mine 444. B) mine 450; C) mine 228; D) Stone hammer, mine 450; E)

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25 731 Stone hammer, mine 228; F) Stone hammer, mine 228; G) Metallurgical furnace IES

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27 732 14 in Valencina de la Concepción (Seville); H) Copper weapons, crucible and nozzle

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30 733 from Valencina de la Concepción (Seville).

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35 735 **Figure 4.** Pb isotopic composition of analyzed samples. A and B show the groups. B, C, D

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37 736 and F show the subgroups. A: Main groups. B: Enlargement of cluster from group 2 to

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40 737 enhance the difference on the potential provenance for mining districts grouped within

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42 738 subgroups 2A, 2B, 2C and 2D. C: Enlargement of cluster from subgroups 2A and 2B

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45 739 to enhance the difference on the potential provenance for mining districts grouped

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47 740 within subgroups 2A1, 2A2, 2B1 and 2B2. In the first diagram above the analytical

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49 741 error is less than the size of the symbols.

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60 746 **Appendix - Table 1 and Table 2**

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1 748 **Table 1:** Samples studied with their geographic location, mine name and coordinates. The

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3 749 assignment to different groups/subgroups is done on the basis of their Pb isotopic

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6 750 composition (see text). (*IPB Ms & Gossan: Iberian Pyrite Belt massive sulphides and

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8 751 gossan).

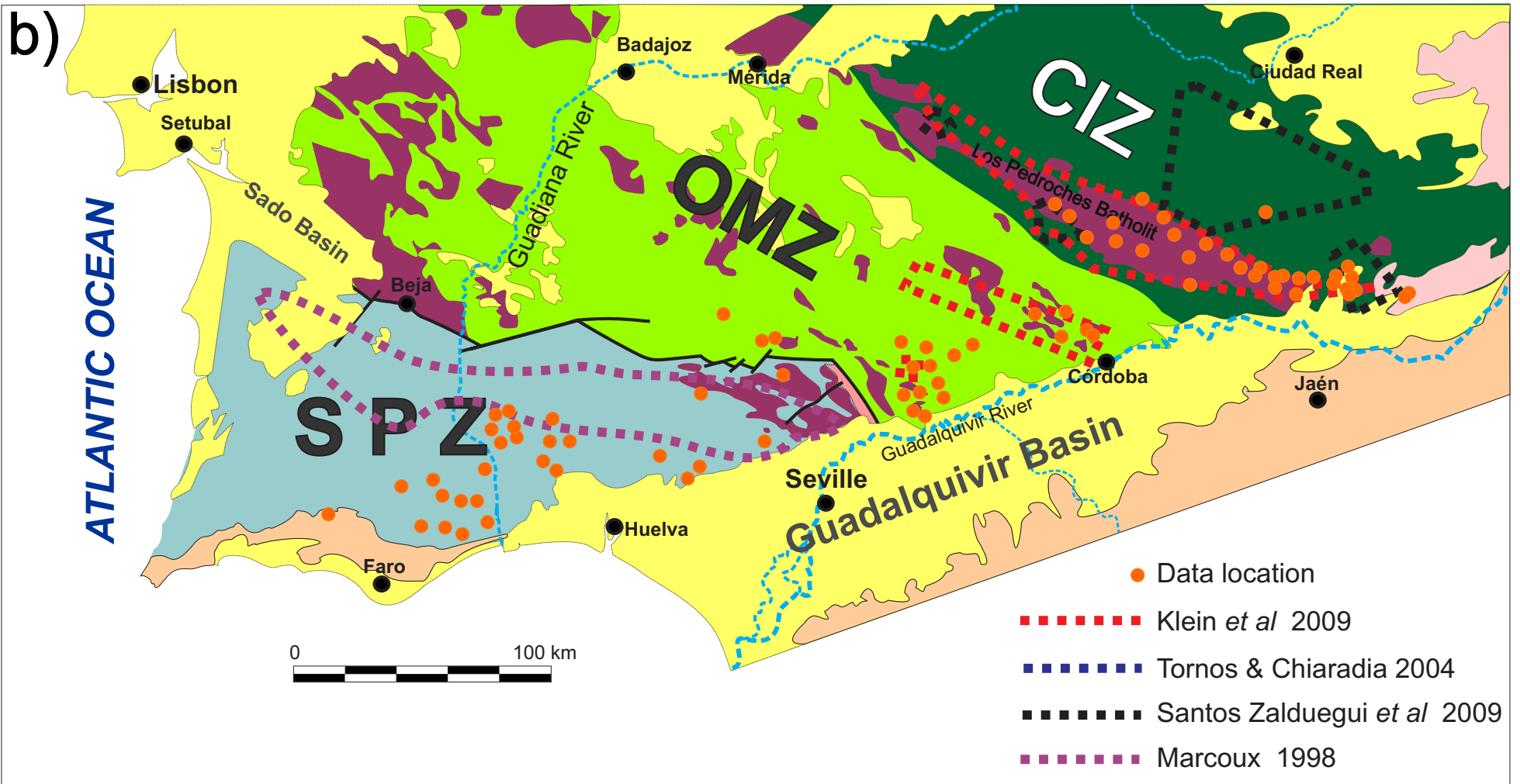
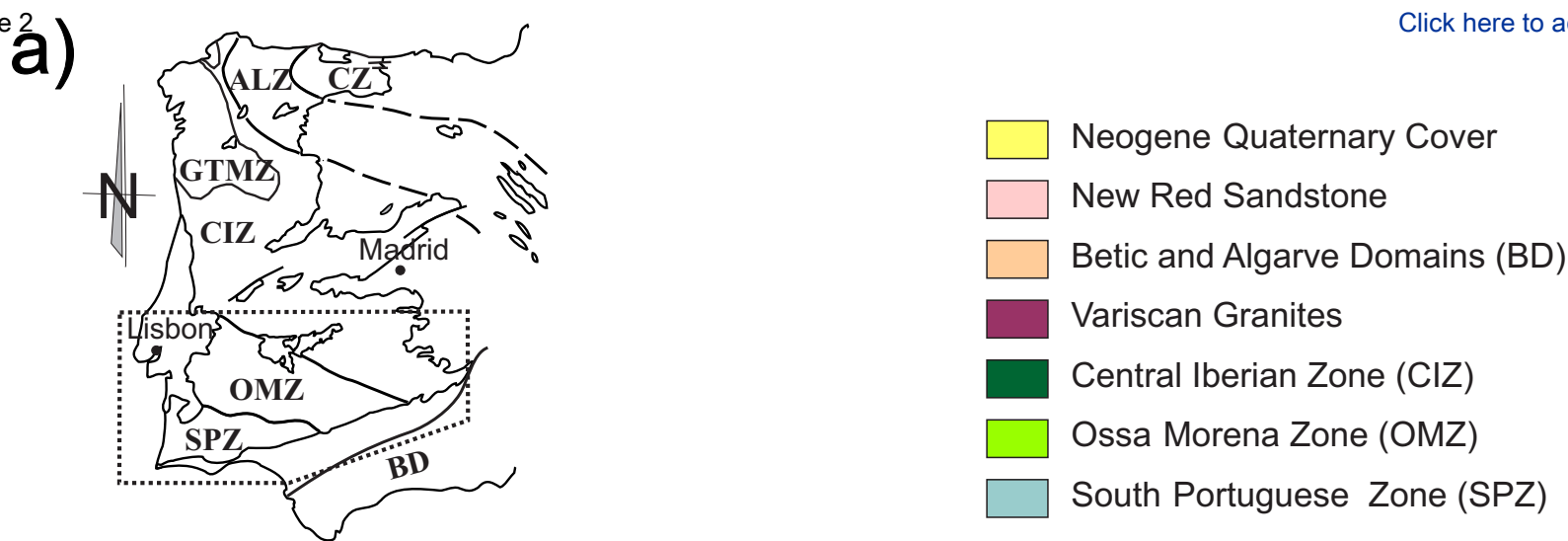
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13 753 **Table 2:** Lead isotope data for the samples studied. The assignment to different

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15 754 groups/subgroups is done on the basis of their isotopic composition (see text).

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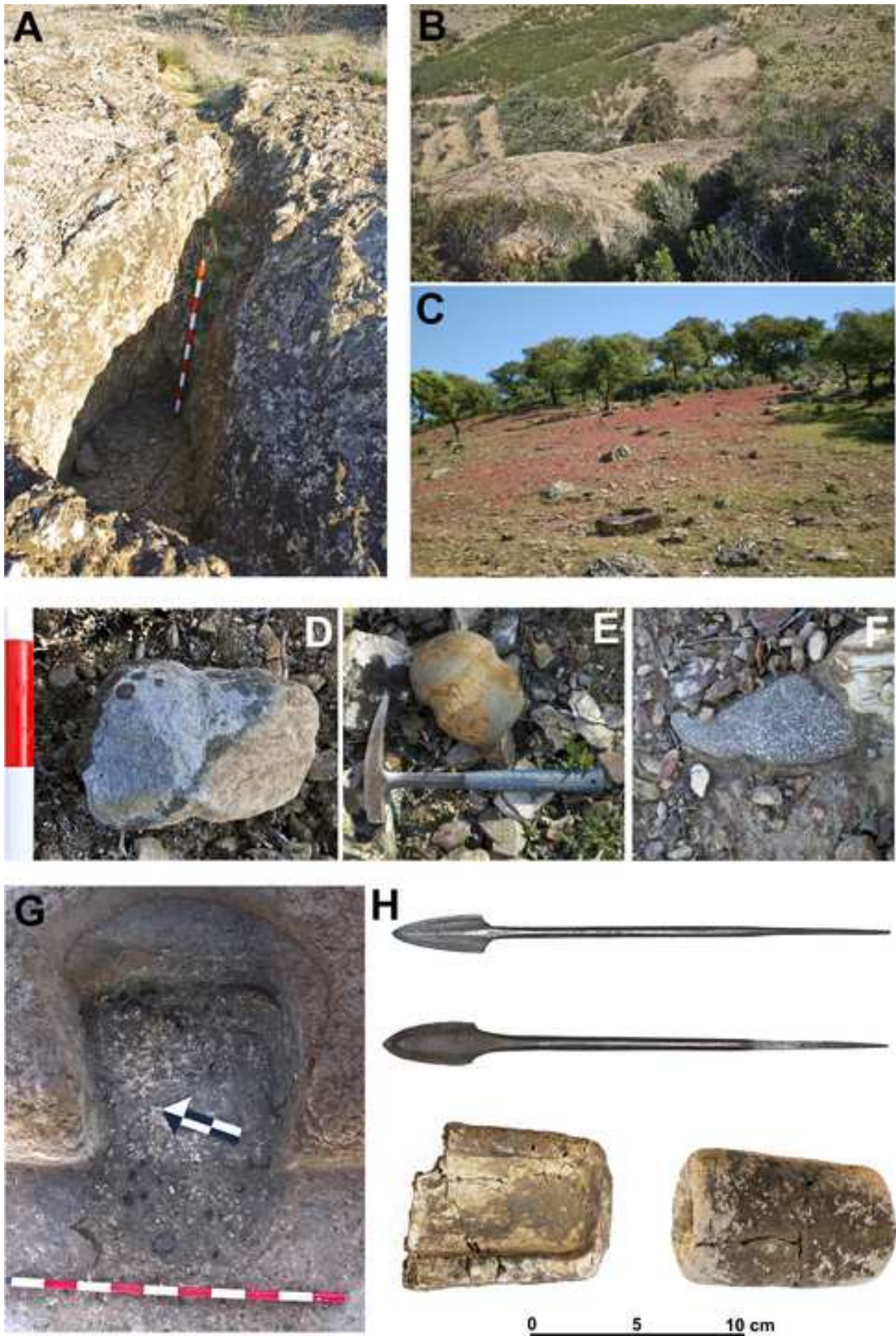
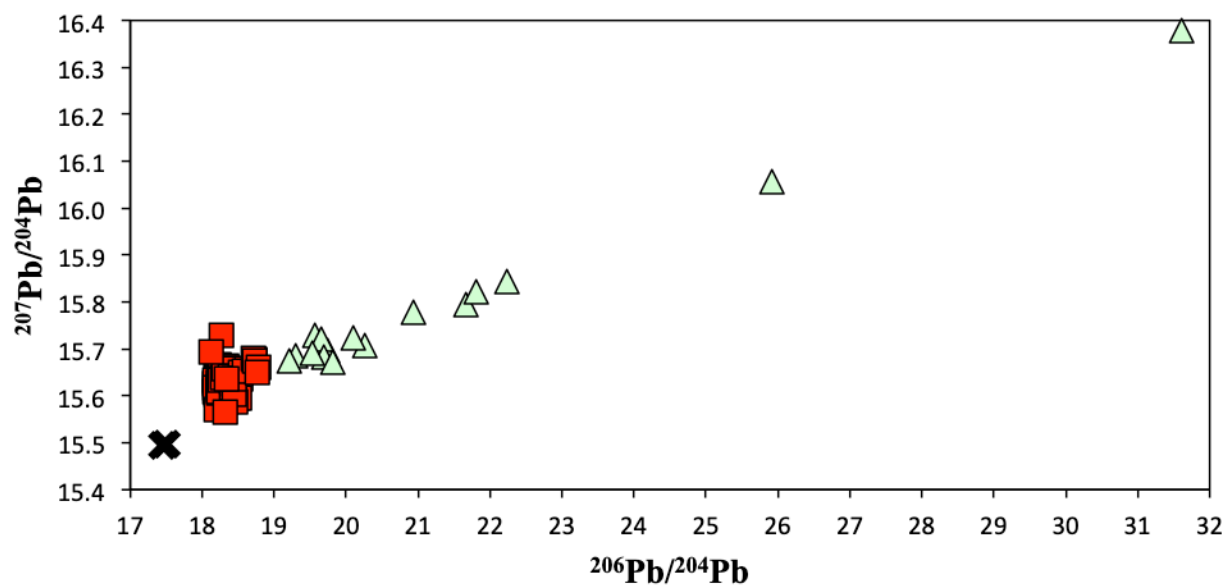


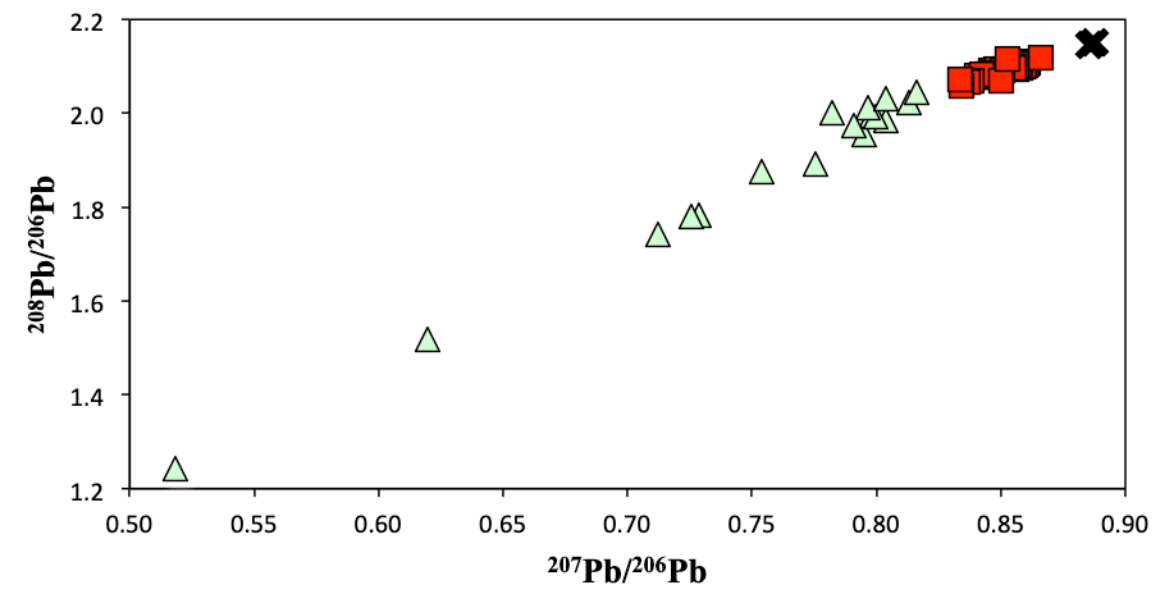
Figure 4

✕ Group 1 ■ Group 2 ▲ Group 3

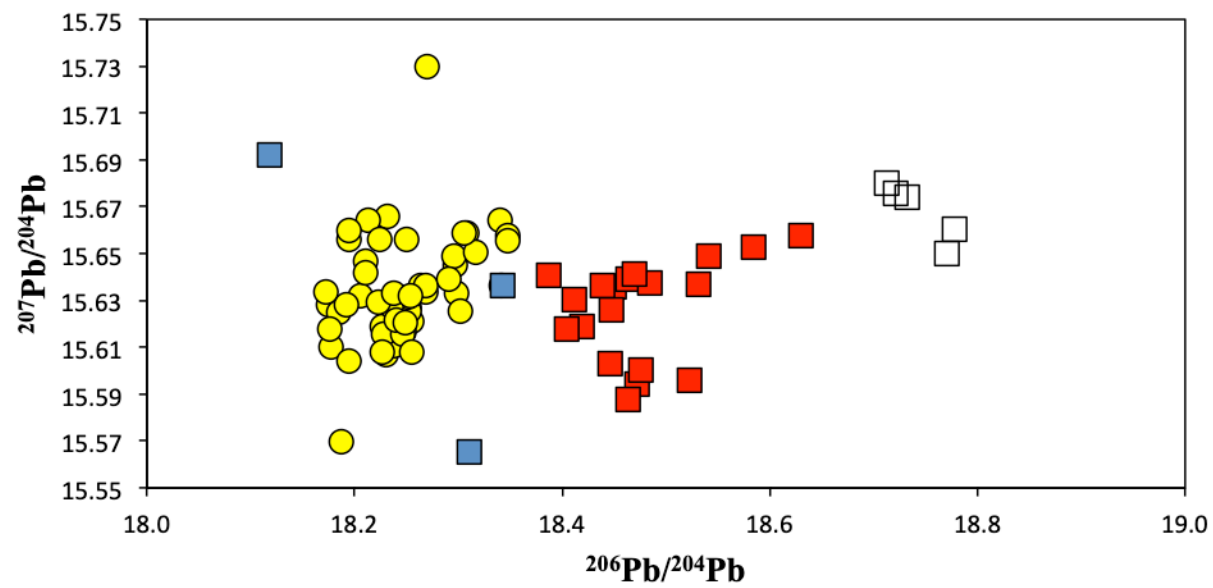


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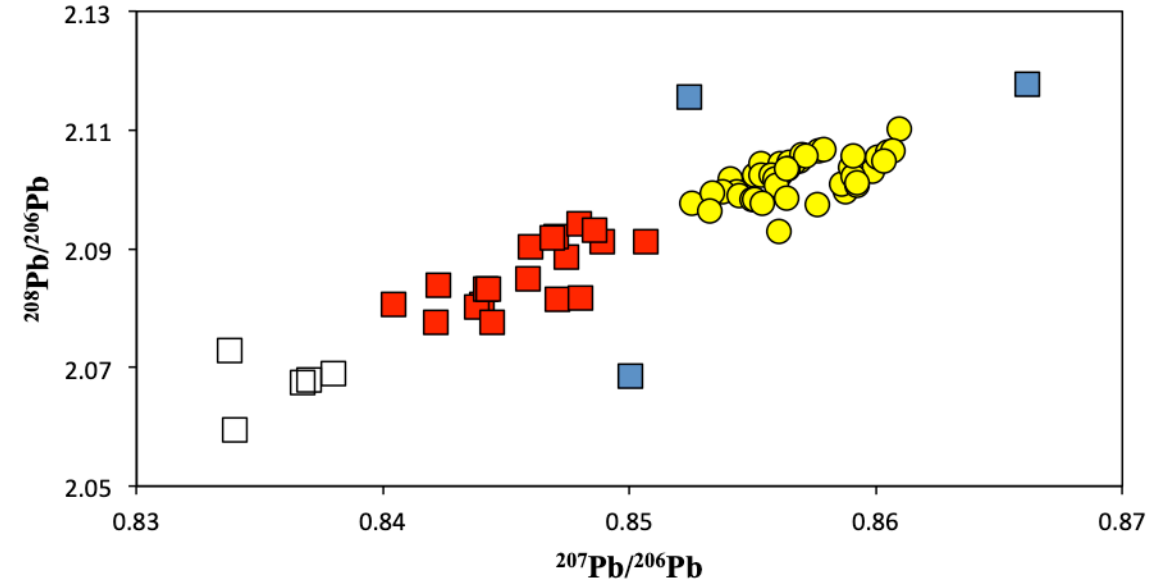
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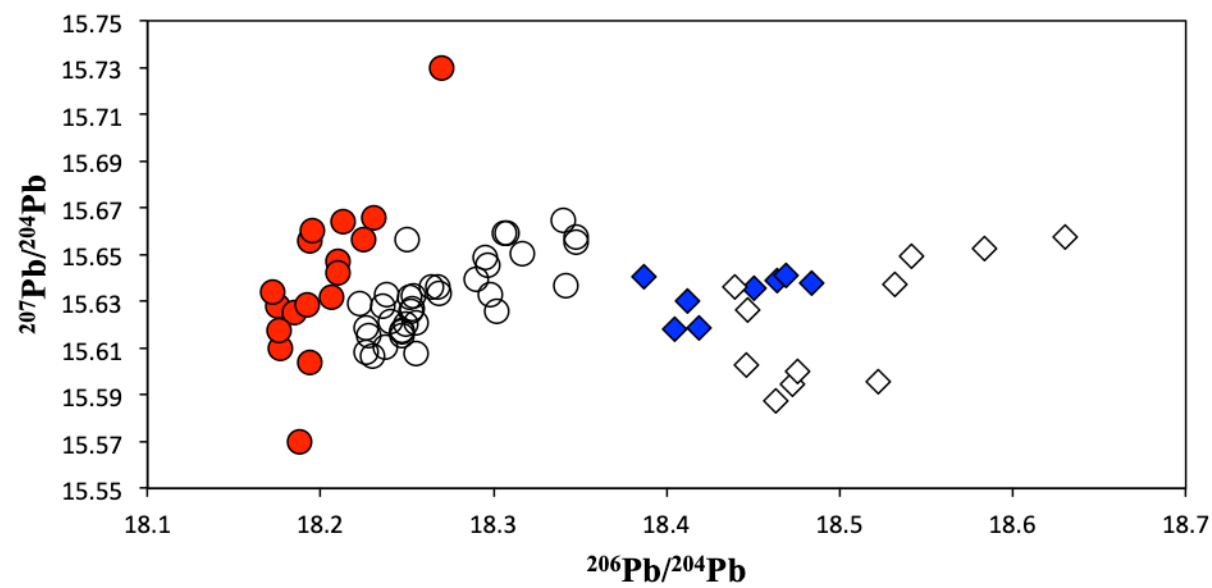
● 2A ■ 2B □ 2C ■ 2D



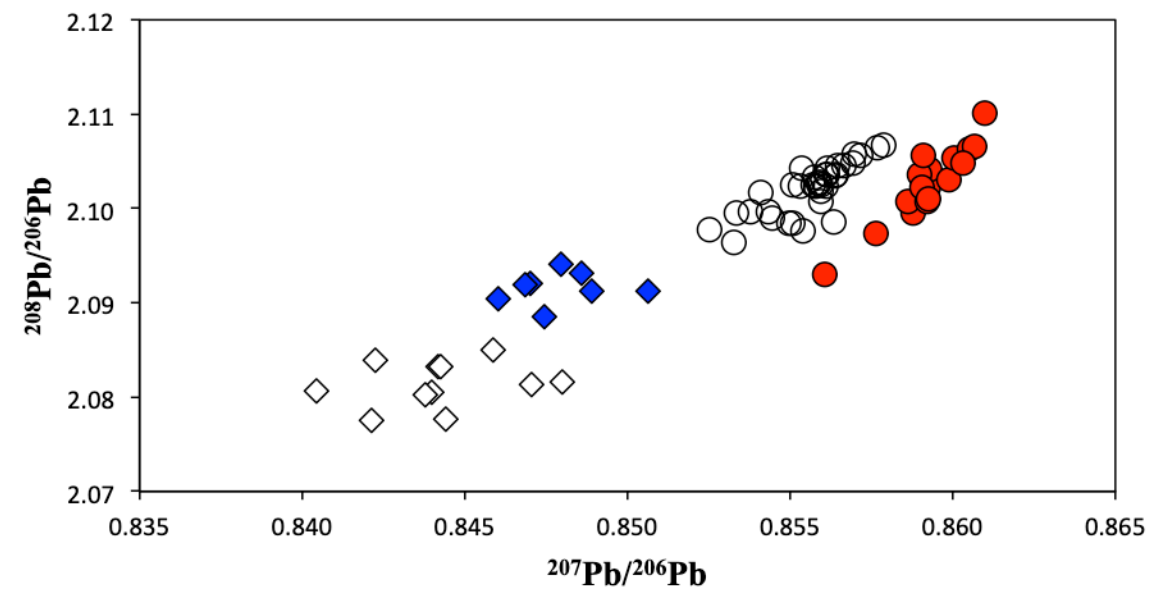
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● 2A1 ○ 2A2 ◆ 2B1 ◇ 2B2

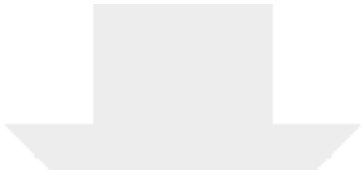


● 2A1 ○ 2A2 ◆ 2B1 ◇ 2B2

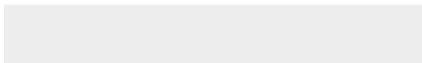
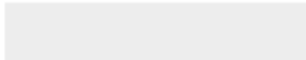





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Supplementary Material
Table 1.docx



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Table 2.docx





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